PALYNOLOGICAL AND PALEOENVIRONMENTAL ANALYSES OF THE MIOCENE SEDIMENTS OF GAP-1 WELL IN THE ONSHORE NIGER DELTA BASIN.

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ABSTRACT

A comprehensive palynological investigation has been carried out for Gap-1 well penetrating the Agbada Formation in the Niger Delta Basin. The identification of shaly-sandstone, shale, mudstone, and clayey-sandstone lithofacies units and the delineation of four (4) informal palynological biozones (GI, GII, GII, and GIV), correlating with the P670, P680, and P720Niger Delta pollen zonations, suggested the Early to Middle Miocene age for the investigated Gap-1 well in the Agbada Formation. The interpretation for the common to the relatively abundant occurrence of land-derived palynomorphs such as Verrucatosporites sp., Laevigatosporites sp, Zonocostites ramonae, Sapotaceae, as well as the occurrence of dinoflagellate cysts, Pediastrum sp, amongst others, as a fluvial-dominated delta to shallow marine environment within a humid climate is confirmed.

INTRODUCTION

The Niger Delta, a prolific hydrocarbon region in West Africa, has been the subject of extensive geological research due to its sedimentary architecture complex and economic significance (Nton & Ogungbemi, 2011; Alege et al., 2020a; Alege, 2022; Odinaka et al., 2024). Palynostratigraphy has emerged as a vital tool in studying the stratigraphy paleoenvironmental and conditions of different world basins by offering insights into the age, depositional environments, and potential hydrocarbon sedimentary reservoirs in sequences. Numerous studies have investigated the palynostratigraphy and paleoenvironmental evolution of the Niger Delta, with works such as that of Evamy et al. (1978) and Jan du Chêne et al. (1978), laying the foundation for biostratigraphy understanding the and stratigraphic frameworks in the delta. These foundational studies established zonal schemes based on palynomorph assemblages, aiding in correlating sedimentary layers across the delta. Similarly, other researchers, such as Alege et al. (2020b) and Alege et al. (2023a), also conducted palynological analyses of the Campanian-MaastrichtianMamu Formation of the Anambra basin identify to its paleoenvironment during the Late Cretaceous period. Aigbadon et al. (2023) employed the palynological, petrological and geochemical attributes of sediments of the Southern Bida Basin to interpret its provenance and paleoenvironment of deposition.

Subsequent research expanded on these early works, with notable contributions by Petters (1995), who examined the paleoenvironmental changes in the Niger Delta using micropaleontological data, and Oloto (1989), who emphasized the role of palynology in sedimentary sequences dating and reconstructing past environments. Furthermore, Osokpor and Oboh-Ikuenobe provided detailed (2008)palynofloral analyses, contributing to understanding the delta's stratigraphic history. However, despite these advancements, gaps remain in the detailed palynostratigraphic zonation and paleoenvironmental interpretation of specific wells, particularly in underexplored onshore areas.

The GAP-1 well, located in the onshore Niger Delta, presented an opportunity to address these research gaps. Previous works have focused mainly on offshore or regional stratigraphic interpretations, leaving the palynostratigraphy of specific onshore wells, such as GAP-1, less understood. Detailed palynostratigraphic and paleoenvironmental analyses of sediments from the GAP-1 well could enhance the biostratigraphic resolution for this region and provide a more refined understanding of the depositional environments that characterised the onshore Niger Delta during different geological periods.

This study aimed at filling the gap by conducting palynostratigraphic a and paleoenvironmental analyses of sediments from the GAP-1 well with the objectives of analysing palynomorph assemblages and identifying age-diagnostic taxa. This research refined the biostratigraphic framework and reconstructed the paleoenvironmental conditions of deposition of these sediments. The findings contributed to a better understanding of the onshore Niger Delta's stratigraphic architecture and its implications hydrocarbon exploration for and environmental changes over geological time.

General Geology of the Area

The Niger Delta Basinis located on the continental margin of West Africa. It covers an area of approximately 75,000 square kilometres, extending from the Gulf of Guinea to onshore Nigeria (Weber & Daukoru, 1975).

The delta developed during the Late Cretaceous to Tertiary periods as a result of the interplay between the tectonic activity associated with the opening of the South Atlantic Ocean and the enormous sediment supply from the Niger River system (Burke, 1972).

The geology of the basin is characterised by a thick sedimentary sequence exceeding 12,000 meters in some areas, consisting of three primary lithostratigraphic units: the Akata Formation, the Agbada Formation, and the Benin Formation (Short &Stauble, 1967). These units reflect a progressive transition from deep marine to coastal and fluvial environments (Figure 1)

The Akata Formation is the deepest unit, consisting of marine shales, silts, and clays. It was formed in deep marine environments and is the primary source of hydrocarbons in the Niger Delta. It ranges in age from Paleocene to Recent. Above the Akata Formation is the Agbada Formation, consisting of alternating sandstones and shales, representing deltaic environments. This formation contains the primary hydrocarbon reservoirs. with sandstone beds serving as reservoir rocks and shale units acting as seals (Stacher, 1995). The Agbada Formation developed during the Eocene to Recent. The Benin Formation is the shallowest and youngest, consisting of continental sands and gravels deposited in a fluvial environment. It is Miocene to Recent in age and is mainly non-hydrocarbon-bearing but is important for groundwater aquifers (Doust&Omatsola, 1990; Reijers, 2011).

The Niger Delta Basin is influenced tectonically by growth faulting and rollover anticlines, which are very significant as hydrocarbon traps (Doust & Omatsola, 1990; Alege, 2017b; Rotimi et al., 2022; Odinaka et al., 2024). The tectonic and sedimentary evolution of the delta, combined with the vast organic-rich sediments deposited in its deep marine environments, have made the basin one of the most prolific petroleum provinces in the world.



Figure 1. The stratigraphic column shows the three lithostratigraphic formations of the Niger Delta (Lawrence et al., 2002).

MATERIALS AND METHODS

Ninety (90) ditch-cutting samples of the GAP-1 well ranging from 3520 to 6190ft and composited at 30ft were processed and analysed for their palynomorphs contents.

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Careful physical examination identified the dominant and minor lithologies, textures, colours, and other accessory mineral components, enabling a detailed lithologic description of the samples (Alege et al., 2015; Idakwo et al., 2013; Alege, 2022; Aigbadon et al., 2023).

A palynological processing method for the maceration of samples modified from Traverse (2007) was applied. This involved crushing samples into fragments, weighing 25 g of each of the samples and alternate usage of inorganic reagents such as dilute Hydrochloric acid (Hcl), concentrated Hydrofluoric acid (HF) and Nitric acid (HNO₃) to macerate the sediments. These reagents removed carbonate and silicate minerals and helped concentrate

the palynomorphs. The acidic reaction was neutralised using potassium hydroxide (KOH) and distilled water. Subsequently, the palynomorph-rich residue was mounted on labelled glass slides using a Norland mounting medium microscopic for analysis. Photomicrographs of the palynomorphs were taken using a camera, and the Strataburg 2.0 software was used in the statistical analysis plot. The results were used to decipher the age and possible paleoenvironment of deposition of the well succession.

RESULTS AND DISCUSSION

Lithofacies

The lithofacies units of Gap-1 have been identified in the ditch-cutting samples by examining the physical and textural characteristics of the sediments (Billman, 1992; Alege, 2017b; Adamu et al., 2018a; Alege et al., 2020a; Oretade and Ali, 2021; Alege et al., 2024; Aigbadon et al., 2024). Four lithofacies units were identified in the study: the shaly sandstone, shale, mudstone, and clayey sandstone units.

Shaly-sandstone facies

These facies represented about 60% of the constituent of the entire Gap-1 well. The facies comprise 85-90% sandstone and 10% shale constituent. The sandstones are mainly fine to medium-grained and sometimes coarse-grained, yellow to brown, and moderately sorted. The minor shale constituents were grey and less fissile. Thesefacies units were well represented at the following depths: 3520-3550 ft, 3760-3880 ft, 4030- 4150 ft, 4600-4630 ft, 5230- 5440 ft and 5860- 6190ft.

Shale facies

The Shale facies represent about 15% of the entire well-depth interval. They are dark grey, soft to moderately hard, and non-fissile. They occurred at depth intervals of 3910-4000ft, 4180-4330ft, 4480- 4950ft, and 4720- 4750ft.

Mudstone facies

The mudstone facies is represented in Gap-1 well by a mixture of shale (30%), clay (40%), and fine to medium-grained sandstone (30%) constituents at depth intervals of 3550 - 3760ft, 4330 - 4390ft, 4720 - 4750ft, 4950 - 5050ft, 5140 - 5230ft, and 5740 - 5860ft. The colours range from brown to grey. This facies represents about 20% of the entire well.

Clayey-sandstone facies

The clayey sandstone facies consist of about 10 - 15% constituent of clay to 85 - 90% sandstone occurring at depth intervals of 5050 - 5140ft and 5440 - 5680ft. The sandstones are brownish, fine to medium-grained and well-sorted.

These facies suggest the fluvial-dominated deltaic facies to shallow marine facies of the transitional paralic sequence of the Agbada formation and are interpreted to be from the fluvial-dominated delta to shallow marine environment (Chukwu, 1991; Reijers, 2011; Nton&Ogungbemi, 2011; Alege, 2017a; Alege et al., 2020a; Alege, 2022).

Palynomorph distribution

Palynomorph analysis was carried out on ninety (90) ditch-cutting samples of the GAP-1 well, ranging from 3520 to 6190 feet. A total of sixty-eight (68) palynomorphs species consisting of forty-seven (47) species of pollens, twelve (12) species of spores, six (6) dinoflagellate spores, diatom frustules, fungal spores and *Pediastrum* were recovered from the well with the percentage proportions of 69.1% (pollens), 17.6% (spores), 8.82% (dinoflagellate cysts), 2.0% (diatom frustules), 2.0% (fungal spores) and 2.0% (freshwater algae, *Pediastrum* sp.) respectively.

The studied interval (3520 to 6190 ft) was characterized by poor to moderately abundant and diverse palynomorphs, which were dominated by terrestrial species such as Laevigatosporites sp., Zonocostites ramonae, Acrostichum aureum, **Psilatricolporit** *Retitricolporites* irregularis, escrassus, *Verrucatosporites* sp., Sapotaceae and catatumbus. *Striatricolporites* Low to common records of fungal spores and freshwater algae Botryococcus braunii also characterised the studied section. Sparse occurrence of dinoflagellate cysts such as Lingulodinium machaerophorum, Spiniferites sp. and Operculodiniumc entrocarpum were also identified (Figure 2).



Figure 2: Photomicrographs of some palynomorph species from GAP-1 well.

(1).Laevigatosporites sp. (3520 ft) (2). Peregrinipollisnigericus (4480 ft) (3). Verrucatosporites sp. (4060 ft) (4). Sapotaceae (4450 ft) (5). Retibrevitricolporitesobodoensis (5020 ft) (6). Spirosyncolpitesbrauni (6070 ft) (7) Botryococcusbraunii (4510 ft) (8). Zonocostitesramonae (4030 ft) (9). Striatricolporitescatatumbus (4000 ft) (10). Pachydermitesdiederixi (6160 ft) (11). Retitricolporitesirregularis (3820 ft) (12). Crassoretitriletesvanraadshoveni (4030 ft) (13). Monoporitesannulatus (5350 ft) (mag. X400)

Palynostratigraphy

Stratigraphic significant species (marker species) and diagnostic palynoflora associations were identified and used to define the ages of the sedimentary successions penetrated by GAP-1. Figure 3 presents the stratigraphic chart of the palynomorphs.



Figure 3. Stratigraphic distribution of palynomorphs in GAP-1 well

Palynostratigraphic zonation

Four (4) informal biozones, namely GI, GII, GIII, and GIV, were established based on the association of recovered marker species and palynomorphs (Figure 3). These zones correlate with the P670, P680, and P720 subzones of Evamy*etal*. (1978). Thus, the Early to Middle Miocene age is interpreted for the studied interval. Detailed results of the palynological zone are presented in Figure.



Figure 4. Summary of the palynological zones identified in the GAP-1 Well (3520-6190 feet)

Zone GIV

Depth: 3520 - 4030 feet

Age: Middle Miocene (Langhian)

The top of this zone coincides with the first sample analysed and, as such. is stratigraphically higher than the first sample at 3520 feet. The base is defined by the base Crassoretitriletes occurrence of vanraadshoveni at 4030 feet. This interval is characterised by the further common occurrence of *Proteacidites* cooksonii,

Magnastriatite showardi, Pachydermit esdiederixi, the rare presence of Belskipolliselegans and abundant Zonocostites ramonae (Figure 4).

In addition, rare occurrences of dinoflagellate cysts such as *Spiniferitessp.* and *Operculodinium centrocarpum* were identified. This interval correlates with the P720 subzones of Evamy *et al.* (1978).

Zone GIII

Depth: 4030 – 4930 feet

Age Middle Miocene (Langhian) – Early Miocene (Burdigalian)

The top occurrence of Striamonocolpites defines the base of this zone rectostriatus at 4930 feet. The top is defined by the base of Crassoretitriletes occurrence vanraadshoveni at 4030 feet. Additional palynoflora events characterising this interval include *Retibrevitricolporites* obodoensis. Pachydermites diederixi, Retitricolporites irregularis and Sapotaceae. A few fungal spores and diatom frustules were also identified (Figure 4&5). This interval correlates with the P680 subzone of Evamy et al. (1978).

Zone GII

Depth: 4930 - 6010 feet

Age: Early Miocene (Burdigalian) – Middle Miocene (Langhian)

The base of this zone is defined by the (?) quantitative base occurrence of Pachydermites diederixiat 6010 feet. In comparison, the top is defined by the top occurrence of Striamonocolpites rectostriatus at 4930 feet. The common occurrence of Zonocostites *Pachydermites* diederixi. ramonae. and Acrostichum aureum with low records of **Retibrevitricolporites** obodoensis also characterised this interval (Figures 4 and 5). Sparse Lingulodinium machaerophorum, *Leoisphaeridia* sp. And **Operculodinium** centrocarpum were also identified. This zone relates to the P680subzone of Evamy et al. (1978).

Zone GI

Depth: 6010 - 6019 feet

Age: Early Miocene (Burdigalian)

The base of this zone coincides with the last sample analysed, while the top is defined by the (?) quantitative base occurrence of *Pachydermites diederixi* at 6010 ft (Figures 4 and 5). Additional palynoflora events include common *Monoporites annulatus, Acrostichum aureum Verrucatosporites* sp., *Sapotaceae* and *Striatricolporites catatumbus*. This zone correlates with the P670 subzone of Evamy et al. (1978).

Paleoenvironment

The paleoenvironmental analysis was based on the different species of palynomorphs within the study (Figures 3 and 4). The relative abundance of land-derived palynomorphs such as *Verrucatosporitessp., Laevigatosporitessp., Acrostichumaureum, Zonocostitesramonae, Sapotaceae, Striatricolporites catatumbus* and *Pachydermites diederixi* amongst others, were interpreted to be predominantly fluvio-deltaic to brackish swamp water environment of deposition within a humid climate (Reijers et al., 1997; Idakwo et al., 2013; Ogbahon, 2019; Aigbadon et al., 2022; Alege et al., 2023).

of dinoflagellate The presence cysts: Operculodinium centrocarpum (4060ft, 5290ft and 5710ft) and Spiniferites sp. (3640ft, 3700ft) with moderate records of pollen, spores and Botryococ cusbraunii indicate a shallow marine environment with frequent freshwater incursions (Wrenn & Kokinos, 1986; Bolaji et al., 2020). Similarly, the lone occurrence of Pediastrumsp (5440ft) and freshwater algae suggests a progradational marginal marine environment (Tahoun et al., 2017; Aigbadon et al., 2023). Diatom frustules in the study also indicate a shallow marine deposition environment (Henchiri, 2007).

Thus, the occurrences of terrestrial palynomorphs and freshwater algae, along with dinoflagellate cysts and diatom frustules, give credence to a fluvial-dominated deltaic environment to a shallow marine environment of deposition within a wet to humid climate (Garzon et al., 2012; Adamu et al., 2018b; Ogbahon, 2019; Alege et al., 2020a).

CONCLUSION

The investigation of palynomorphs recovered from ninety (90) ditch cuttings obtained from Gap-1 well has aided in interpreting the palynostratigraphy and paleoenvironment in the onshore Niger Delta basin. Four lithofacies units were identified in the study: the shaly sandstone, shale, mudstone, and clayey sandstone units. These facies suggested the fluvial-dominated deltaic facies to shallow marine facies of the transitional paralic sequence of the Agbada formation. They were interpreted to be from the fluvial-dominated delta to shallow marine environment.

Four (4) informal biozones, namely GI, GII, GIII, and GIV, were established based on the association of recovered marker species and palynomorphs. These zones correlated with the P670, P680, and P720 subzones of Evamyetal. (1978). Thus, the Early to Middle Miocene age are interpreted for the studied interval. The paleoenvironmental analysis was based on the different species of palynomorphs within the study. The common occurrence of land-derived palynomorphs Verrucatosporites such as sp., Laevigatosporites sp, Zonocostites ramonae, Sapotaceae, Striatricolporites catatumbus and Pachydermites diederixi amongst others, are interpreted to be predominantly fluvial-deltaic to brackish swamp water environment of deposition within a humid climate.

occurrences of terrestrial Thus. the palynomorphs freshwater and algae (Pediastrum sp.) along with dinoflagellate (Operculodinium cysts centrocarpum, Spiniferites sp.) and diatom frustules, give credence to a fluvial-dominated deltaic environment to a shallow marine environment of deposition in a humid climate in the Agbada Formation of the Niger Delta.

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